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Rock mass statistical homogeneity investigation along a highway corridor in Vietnam

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1 **Rock mass statistical homogeneity investigation along a highway** 2 **corridor in Vietnam**

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12

13 **Abstract**

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15 or rock mechanics literature to firmly demarcate between the statistical homogeneity and non-homogeneity
16 of rock masses. The study shows that the available methods can only be used to rank rock masses with
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18 comparison, (b) linear correlation coefficient analysis and (c) contingency table analysis, are applied in this
19 paper to evaluate the statistical homogeneity of rock masses that exist along a 15 km stretch of the 3B
20 Highway corridor in Vietnam. The study clearly illustrates that application of different methodologies can
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22 difficulty in making a judgement on rock mass statistical homogeneity using results obtained from only one
23 method. The paper shows how better decisions were made on rock mass statistical homogeneity by
24 comparing the results obtained from the three different methodologies using only discontinuity orientation
25 data.

26 **Keywords:** Rock discontinuity orientation, Rock mass statistical homogeneity, polar equal area pole
27 plots, linear correlation analysis, contingency table analysis.

28

29 **Declarations**

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35

36 *Conflict of interest/competing interest*

37 The authors declare that they have no conflict of interest.

38

39 *Availability of data and material*

40 Possibility exists to make data available after publishing the papers in journals.

41

42 *Code availability*

43 Not applicable

44

45 **1. Introduction and literature review**

46 The mechanical and hydraulic behaviour of discontinuous rock masses as well as rock slope stability
47 depend very much on the discontinuity geometry pattern of the rock masses ([Kulatilake et al. 1993 and](#)
48 [1996](#)). The discontinuity geometry pattern in a rock mass can vary from one region to another. The first
49 step in determining the discontinuity geometry pattern in a rock mass is the identification of statistically
50 homogeneous regions ([Kulatilake et al. 1993 and 1996](#)). It is also known as determination of structural
51 domains in the traditional rock mechanics literature. Once the statistically homogeneous regions are
52 determined, each one of these regions may be represented by a discontinuity geometry model. Such a
53 discontinuity geometry model can then be used to study the mechanical or hydraulic behaviour or the rock
54 slope stability of the rock mass. For complete statistical homogeneity of rock mass regions, the discontinuity
55 sets should have similar distributions for discontinuity intensity in three dimensions, orientation, spacing,
56 size, shape, roughness, and discontinuity constitutive properties. However, it is very difficult or almost
57 impossible to achieve the statistical homogeneity to this level in rock engineering practice. At present, in
58 practice, only the number of discontinuity sets and their orientation distributions are considered in
59 determining the statistical homogeneity of rock masses. In practice, mostly, such rock masses are
60 determined by visually comparing samples of geologic structure orientations, each of which consists of
61 discontinuity normal vectors (also known as poles) plotted on a polar equal-area net. When discontinuity
62 orientations do not show definite pole clusters, visual comparisons are often not sufficient to evaluate
63 statistical homogeneity between chosen regions. Even when discontinuity orientations show definite pole
64 clusters, in many cases it is difficult to compare the distribution of poles for a discontinuity set in the

65 considered two regions only through visual inspection. In such situations, it may be desirable to use the
66 statistical procedures suggested by [Miller \(1983\)](#), [Mahtab and Yegulalp \(1984\)](#), [Kulatilake et al. \(1990\)](#)
67 [and Martin and Tannant \(2004\)](#), and the fractal procedure suggested by [Kulatilake et al. \(1997\)](#), in
68 addition to the visual examination of pole plots plotted on a polar equal-area net to investigate statistically
69 homogeneous regions in rock masses. Attempts have been made to investigate statistical homogeneity of
70 rock masses in a few sites in the world using discontinuity geometry data ([Kulatilake et al. 1990, 1996](#)
71 [and 1997](#); [Martin and Tannant 2004](#); [Phi et al. 2012](#); [Thanh et al. 2015](#); [Li et al. 2015](#); [Song et al.](#)
72 [2018](#)). In these papers' authors had to make subjective decisions on statistical homogeneity of rock masses
73 under significant uncertainty. The reason for that is the lack of a solid objective procedure in the current
74 rock mechanics or engineering geology literature to firmly demarcate between the statistical homogeneity
75 and non-homogeneity for real world rock masses. It is important to mention that decision making on
76 statistical homogeneity for real world sites is a very difficult, challenging assignment. It is not a trivial task
77 at all. It cannot be made in a firm, deterministic fashion. However, this aspect is not clearly shown through
78 detailed case studies in the literature. One of the main aims of this paper is to show this difficult and
79 challenging aspect through a very carefully performed detailed case study. Only joint orientation values are
80 required to obtain pole plots on a polar equal area net, and to apply [Miller \(1983\)](#), [Kulatilake et al. \(1990\)](#)
81 [and Martin and Tannant \(2004\)](#) procedures. On the other hand, discontinuity trace maps are required to
82 apply the fractal procedure given by [Kulatilake et al. \(1997\)](#). This paper first reviews the salient features
83 and the capabilities of the polar equal area pole plots and [Miller \(1983\)](#), [Kulatilake et al. \(1990\)](#) and
84 [Martin and Tannant \(2004\)](#) procedures. The paper then shows how modified versions of these procedures
85 are applied to 4189 orientation data to investigate statistical homogeneity of rock masses located along a
86 highway corridor in Vietnam. The obtained results very clearly show the difficult and challenging nature
87 of decision making on statistical homogeneity of rock masses.

88

89 **Fig. 1.** Trajectories of the maximum compressive stress within the Indochina Peninsula during (a) the
90 Oligocene and (b) at the present time ([Kasatkin et al., 2014](#)). The legend on the map in Figure 1 is described
91 as follows: 1. trajectories of the maximum compressive stress: (a) which are directly related to the Indo-
92 Eurasian plate collision and (b) its far-field effects; (2) faults and directions of displacement (arrows); (3)
93 zone of continental collision; (4) subduction zone; (5) extension structures; (6) spreading zones; (7) current
94 position of the land; Red River Fault System (RRFS); Cao Bang - Tien Yen fault (CB - TY).

95

96 **2. Geological conditions of the considered site**

97 The study area lies between two major fault zones (Red River fault zone and Cao Bang- Tien Yen fault
98 zone) with the faults striking in the NW-SE direction in the Xuat Hoa area, Bac Kan Province (northeast
99 region) of Vietnam (**Fig. 1**). Earlier studies ([Nguyen, 1991](#), [Phung et al. 1996](#)) suggested that the study
100 area has undergone two major phases of tectonic activity; an early phase compressive tectonic activity

101 occurred in the E-W direction starting from the Eocene and extending to the Oligocene-Miocene period
102 followed by a late phase compressive tectonic activity in the N-S direction beginning from the Pliocene to
103 the present (**Fig. 1**). Based on results of an extensive analysis of fault plane striations, **Thanh (2019)**
104 identified four main compressive tectonic activity phases in the study area; the order and direction of
105 occurrence of these phases are (a) NW-SE, (b) E-W, (c) NE-SW, and d) N-S. The geological map of the
106 study area (**Nguyen et al. 2000**) along with the survey stations (SS) used to collect discontinuity orientation
107 data are shown in **Fig. 2**. Note that the survey stations are selected along a highway named as the 3B
108 Highway. As shown in **Fig. 2**, it is a highly winding road. It is very important to keep the rock slopes stable
109 along this highway corridor. The rock mass statistical homogeneity of the highway corridor is investigated
110 in this paper as the first step towards investigating rock slope stability. As explained above, the tectonic
111 condition of the study area is quite complex. Therefore, intuitively, a quite complex rock mass statistical
112 homogeneity can be expected along the highway corridor. The dominant lithologies in the study area are
113 Devonian sedimentary rocks. These rocks belong to three main formations as shown in **Fig. 2**: a) Tam Hoa
114 formation (D_{2-3th}) consisting of polymictic conglomerate, gritstone, clayey shale and limestone; b) Mia Le
115 Formation (D_{1ml1} and D_{1ml2}) consisting of clayish siltstone and marlaceous shale; and c) Na Quan
116 formation (D_{1-2nq1} and D_{1-2nq2}) consisting of marlaceous shale and shale interbedded with limestone
117 respectively. Typical rock masses of each of these formations located at a few survey stations are shown in
118 **Fig. 3**.

119

120 **Fig. 2.** Geological map of the study area showing the survey stations (SS) used to obtain orientation data.

121

122 **Fig. 3.** Typical rock masses that exist at a few survey stations.

123

124 Discontinuity orientation data were collected from three formations at 32 survey stations located along
125 a 15 km stretch of the 3B Highway corridor. The location map of the survey stations and the lithologies of
126 the stations are shown in **Fig. 2**. The survey station SS-1, SS-4 through SS-11, and SS-13 through SS-32
127 are from the D_{1ml2} formation. The survey station SS-2 is from the D_{1ml1} formation. The survey station SS-
128 3 is located at the boundary between the D_{2-3th} and D_{1-2nq2} formations. The survey station SS-12 is from
129 the D_{1-2nq2} formation. The survey stations SS-12 and SS-13 are on the opposite sides of a fault line. A very
130 good chance exists to have poor statistical homogeneity when rock masses from opposite sides of a fault
131 line are compared. If the formations are quite different, then there is a tendency to have a low level of
132 statistical homogeneity among the rock masses located in different formations. A total of 4189 discontinuity
133 orientation data were collected from the 3B Highway corridor as given in **Table 1**. A few details about each

134 survey station and the number of discontinuity orientation data collected from each station are presented in
135 **Table 1.**

136

137 **Table 1** Survey station locations, geologic age and the number of discontinuity orientation data collected
138 for the various stations.

139

140 **3. Main features of the procedures used to evaluate statistical homogeneity**

141 *3.1. Polar equal area pole plot comparison*

142 Polar equal area net is used to plot the discontinuity normal vectors (also known as poles) of orientation
143 data and to find the distribution of discontinuity poles for each discontinuity set belonging to a particular
144 rock mass site ([Goodman 1976](#); [Hoek and Bray 2004](#)). The pole plotting can be done either using the
145 lower hemispherical projection or the upper hemispherical projection ([Goodman 1976](#)). In this study the
146 lower hemispherical projection was used in plotting the discontinuity poles. These plots show pole
147 concentrations. These pole concentrations can be used to determine the number of discontinuity sets that
148 exists in a rock mass site. In addition, these pole plots show the distribution of orientation data for each
149 discontinuity set. The polar equal area pole plots of two rock mass sites can be compared visually to assess
150 the statistical similarity of the two sites based on the number of discontinuity sets that exists in the two sites
151 and their orientation distributions. This statistical similarity is also worded as the rock mass statistical
152 homogeneity of the two sites in this paper.

153 In determining the rock mass statistical homogeneity between two sites, efforts were made to draw
154 visual conclusions about the strength of statistical homogeneity of two adjacent regions. This was achieved
155 by using only the polar equal area pole plots to assign qualitative ranks of the rock mass statistical
156 homogeneity. In investigating the statistical homogeneity between two rock mass sites, it is necessary to
157 first check whether the same number of discontinuity sets appear in the two sites. If so, then it is necessary
158 to check for each of the discontinuity sets, orientation wise, whether the same discontinuity set appears in
159 the two data sites being compared. If the same discontinuity sets exist in the data sites, a further check on
160 the similarity of their orientation distributions is conducted.

161 Generally, data sites which are in proximity in the same rock formation or similar rock formations tend
162 to have stronger statistical homogeneity compared to sites that are far apart in the same rock formation or
163 in different rock formations. First, the polar equal-area pole plots were compared among every two adjacent
164 stations along the highway corridor starting from SS-1 versus SS-2 to SS-31 versus SS-32. If any two
165 adjacent stations were found to be strongly statistically homogeneous, then the orientation data of those
166 two stations were combined to make a new polar equal area pole plot for each of the combined stations.

167 Then the polar equal area pole plot of each of the combined stations was compared with the next adjacent
 168 station to evaluate the statistical homogeneity among those stations. This process was continued for all the
 169 adjacent stations along the highway corridor, which showed strong statistical homogeneity.

170

171 *3.2. Linear correlation coefficient analysis*

172 It is often difficult to evaluate the statistical homogeneity only by visually comparing the polar equal
 173 area plots of different sites when definite pole clusters do not appear through the discontinuity orientation
 174 data. Even when the pole clusters are visible, the qualitative nature of the polar equal area pole plot
 175 comparison procedure makes the result subjective. Such difficulties can be avoided by adopting quantitative
 176 methods such as the linear correlation coefficient, ρ , analysis and/or contingency table analysis in
 177 identifying statistically homogeneous domains. This section covers the linear correlation coefficient
 178 analysis. Section 3.3 covers the contingency table analysis. A polar equal area patch network was
 179 constructed by dividing the horizontal circumferential 0 to 360 degrees into 10-degree cells and horizontal
 180 radial 0 to 90 degrees on the polar equal area net into 10-degree cells. This created a 324-patch polar equal
 181 area net as shown in **Fig. 4**.

182

183 **Fig. 4.** The constructed 324-patch polar equal area net.

184

185 Poles of the discontinuity orientation data of each of the 32 stations were plotted using the lower
 186 hemispherical projection on the 324-patch polar equal area net. For each of the two station combinations
 187 considered under polar equal area plot comparison method (see section 3.1), the linear correlation
 188 coefficient, $R(x, y)$, was calculated according to equation (1). The numerator of Eq. (1) is the covariance.
 189 The two square root terms appearing in the denominator of Eq. (1) are the standard deviations of x and y .

190
$$R(x, y) = \frac{\frac{1}{(n-1)} \sum_{i=1}^n (x_i - \bar{X})(y_i - \bar{Y})}{\sqrt{\frac{1}{(n-1)} \sum_{i=1}^n (x_i - \bar{X})^2} \sqrt{\frac{1}{(n-1)} \sum_{i=1}^n (y_i - \bar{Y})^2}} \quad (1)$$

191 In Eq. (1), $(x_1, y_1), (x_2, y_2), \dots, (x_n, y_n)$ are n pairs of pole number observations coming from n 10 by 10
 192 degree cells for each of the considered two stations. \bar{X} and \bar{Y} are the average pole numbers for a 10 by 10-
 193 degree cell from each of the two stations considered in the computation. x_i is the pole number that appears
 194 in cell i of the first station and y_i is the pole number that appears in cell i of the second station. Because a
 195 324-patch polar equal area net is used for the calculations, $n = 324$ for the conducted study. The sorting of
 196 discontinuity poles was done in such a way that no single discontinuity pole falls into two cells in a polar
 197 equal area net, which will result in double counting. R expresses the strength of the similarity or statistical

198 homogeneity of the pole distributions among the considered two stations. R values range from -1 to +1 and
199 do not depend on the discontinuity data numbers of the considered two stations. The R value of +1 implies
200 perfect positive linear correlation while the R value of -1 implies perfect negative linear correlation. The R
201 value of 0 implies no correlation. Low R values imply weak linear correlation and high R values imply
202 strong linear correlation. Higher the magnitude of the R, better the linear correlation and better the chance
203 for the statistical homogeneity. Unfortunately, it is not possible to specify an objective R value which
204 demarcates between the statistical homogeneity and non-homogeneity. Therefore, in this paper, the
205 computed R value is used to assign a rank to the tested two regions to indicate the strength of the rock mass
206 statistical homogeneity.

207

208 *3.3. Contingency table analysis*

209 Application of contingency table analysis to polar equal area plots requires the division of the lower
210 hemisphere into patches that have equal areas either on the lower or upper hemisphere. In this paper, the
211 lower hemisphere is chosen. In this study, a modified version of **Miller's method (1983)** reported by
212 **Kulatilake et al. (1990)** was used to perform contingency table analysis. Based on the modified Miller's
213 method, a formulation is given to divide the lower hemisphere polar equal area net into equal surface area
214 patches after selecting the number of dip bands and the number of patches (cells) in each band. Most of the
215 discontinuities of the survey stations were found to be of sub-verticals giving high dip angles. Therefore,
216 most of the poles were expected to get plotted in the polar equal area net within 0-40 degrees from the
217 periphery towards the center of the patch network. That means it is better to have 4 bands rather than 3,
218 which is usually used, for the patch network and more patches or cells in the outer two bands for better
219 discrimination of pole sorting to perform contingency table analysis. Based on these concepts, for the first
220 time a 68-cell patch network as shown in **Fig. 5** was designed to perform contingency table analysis.

221

222 **Fig. 5.** The designed 68-cell patch network used for the contingency table analysis.

223

224 **Table 2** Contingency table for analyzing the polar equal area plots of different stations (**after Miller, 1983**).

225

226 **Table 2** provides a contingency table for analyzing the polar equal area plots of different stations. Entry
227 of each patch (cell) in the table is the number of discontinuity poles that occur in that patch (f_{ij}). To test the
228 null hypothesis that the rock mass stations are statistically homogeneous, the following chi-square statistic
229 is calculated to perform the chi-square test.

230

$$\chi^2 = \sum_{i=1}^r \sum_{j=1}^r \frac{(f_{ij} - e_{ij})^2}{e_{ij}} \quad (2)$$

231 In Eq. (2), e_{ij} is the expected number of discontinuity poles in the ij cell and it is calculated through $e_{ij} =$
 232 $(R_i C_j) / N$ (Miller 1983; Kulatilake et al. 1990). The symbols R_i and C_j are explained in **Table 2**. Under the
 233 null hypothesis, to have a perfect fit, all f_{ij} should be equal to e_{ij} . In such a case $\chi^2 = 0$. Thus, large values
 234 of χ^2 tend to discredit the null hypothesis, and small values of χ^2 tend to confirm the hypothesis. If the
 235 null hypothesis is true, then in general the calculated χ^2 is chi-square distributed with $(r-1)(c-1)$ degrees
 236 of freedom. According to the usual chi-square test procedure, the null hypothesis is rejected if the χ^2 from
 237 Eq. (2) exceeds the value from the chi-square table corresponding to $(r-1)(c-1)$ degrees of freedom at the
 238 chosen significance level of α . In statistical applications, the statisticians commonly use 0.05 or 0.10 for α .
 239 Another possibility is to obtain the α value from the chi-square table corresponding to the calculated χ^2
 240 value from Eq. (2) using the degrees of freedom $(r-1)(c-1)$. This value, termed as p , can be considered as
 241 one's confidence in accepting the null hypothesis and seems more appropriate to use in engineering
 242 applications.

243 The accuracy of the chi-square test depends on the e_{ij} distribution in the patch network. Miller
 244 recommends use of Lancaster's (Lancaster, 1969) criterion for contingency table analyses of equal area
 245 plots. To satisfy Lancaster's criterion, it will be necessary to try out different patch networks. It is possible
 246 to find equal area plots for which Lancaster's criterion cannot be met. For such situations, p can be evaluated
 247 using the approximation that χ^2 follows a normal distribution under the condition of degrees of freedom
 248 greater than 30.

249 The value obtained for p depends on the following factors: (a) the number of patches in the patch
 250 network; (b) the relative position of the patch network with respect to north; and (c) whether the Lancaster's
 251 criterion was met or not. Therefore, for the same data, it is possible to get several different values for p .
 252 Then one faces the almost impossible task of deciding on a value for p which demarcates between the
 253 statistical homogeneity and non-homogeneity. Therefore, it is not possible to determine rock mass statistical
 254 homogeneity uniquely through this method too as in the previous two methods. However, if the afore-
 255 mentioned three factors are kept the same in testing different regions of the rock mass, then the obtained p
 256 values can be used to rank the tested regions with respect to the strength of homogeneity. The chance for
 257 statistical homogeneity increases with increasing p . In this paper, 18 contingency test results are obtained
 258 for each tested two regions starting from the patch network aligned to north and then by rotating the patch
 259 network using an increment of 10 degrees. For each of these test results χ^2 and p values are obtained based

260 on both the chi-square and normal distributions. These 18 sets of values were then used to compute average
261 χ^2 and average p values based on both the chi-square and normal distributions. These average p values are
262 then used to rank each of the tested two regions with respect to the strength of statistical homogeneity.

263

264 4. Results

265 4.1. Use of polar equal area pole plot comparison to assign qualitative ranks of statistical homogeneity

266 Discontinuity orientation data collected from the 32 survey stations were plotted on the equal area polar
267 net using the Dips 7.0 software package (Rocscience Inc., 2019) and selecting the lower hemisphere
268 projection. The obtained lower hemispherical polar equal area pole plots are shown in **Fig. 6**. These plots
269 show about two sub-vertical discontinuity sets. In certain stations both sets appear prominent. In the
270 remaining stations only one prominent discontinuity set appears. For most of the stations, the orientation
271 distributions and mean orientation directions of these two discontinuity sets vary from one station to another
272 significantly. That weakens the strength of the statistical homogeneity among many stations. By looking at
273 the polar equal area pole plots it is difficult to firmly demarcate between the statistical homogeneity and
274 non-homogeneity between different stations. Therefore, an attempt was made to qualitatively rank the
275 stations with respect to the strength of statistical homogeneity.

276

277 **Fig. 6.** Lower hemispherical polar equal area pole plots of the 32 stations located along the highway
278 corridor.

279

280 The polar equal area pole concentration plots shown in **Fig. 6** were used in making a judgement about
281 the strength of statistical homogeneity for each of the two adjacent stations located along the highway
282 corridor. For each case, a qualitative rank was assigned based on the procedure mentioned under section
283 3.1. Poor (P), Low (L), low to poor (L-P), medium (M), or high (H) ranks were assigned to the cases to
284 indicate the observed visual strength of statistical homogeneity. Note that this ranking procedure is
285 subjective. The obtained results are given in **Table 3**. The stations SS-12 and SS-13 showed a very poor
286 level of statistical homogeneity. Note that these two stations are located on the opposite sides of a fault.
287 This result was intuitively expected. Rock formation differences among the tested two stations exist for the
288 following cases: SS-1 versus SS-2, SS-2 versus SS-3, SS-3 versus SS-4 and SS-11 and SS-12. Note that all
289 those cases show either a P, L-P or L level of statistical homogeneity. Possibility of a such a result was
290 stated in section 3.1.

291

292 **Table 3** Summary results for strength of statistical homogeneity based on different methods.

293

294 The obtained results from the polar equal-area pole plot comparisons of the discontinuity data (**Table**
295 **3**) indicate that the stations SS-20 versus SS-21, SS-26 versus SS-27, SS-27 versus SS-28 and SS-31 versus
296 SS-32 have H level of statistical homogeneity among all the cases tested. These cases show existence of
297 similar discontinuity sets with quite similar orientation distributions. The following cases showed M level
298 statistical homogeneity based on existence of similar discontinuity sets and fairly similar orientation
299 distributions: SS-5 versus SS-6, SS-10 versus SS-11, SS-14 versus SS-15, SS-15 versus SS-16, SS-18
300 versus SS-19, SS-19 versus SS-20, SS-28 versus SS-29 and SS-30 versus SS-31. Orientation data of
301 locations that indicated existence of either H level or M level statistical homogeneity between adjacent
302 stations were then combined as follows to make comparisons among equal area pole plots of combined
303 stations: (a) SS-14 & 15 versus SS-16, (b) SS-18 versus SS-19 & 20, (c) SS-18 & 19 versus SS-20 & 21,
304 (d) SS-26 & 27 versus SS-28, (e) SS-26 & 27 versus SS-28 & 29 and (f) SS-30 versus SS-31 & 32. Lower
305 hemispherical polar equal area pole plot comparisons made for the combined stations are shown in **Fig. 7**.
306 The qualitative rankings made through visual comparison are given in **Table 3**. These rankings were
307 assigned based on the rankings made earlier using P through H levels. Out of all the stations compared, the
308 stations 18 through 21 seem to indicate the highest level of statistical homogeneity.

309
310 **Fig. 7.** Lower hemispherical polar equal area pole plot comparisons made for the combined stations: (a)
311 SS-14 & 15 versus SS-16 (L level), (b) SS-18 versus SS-19 & 20 (H level), (c) SS-18 & 19 versus SS-20
312 & 21 (M level), (d) SS-26 & 27 versus SS-28 (L level), (e) SS-26 & 27 versus SS-28 & 29 (L level) and (f)
313 SS-30 versus SS-31 & 32 (L level).

314
315 *4.2. Use of linear correlation coefficient analysis to assign quantitative ranks of statistical homogeneity*

316 Using the procedure explained in section 3.2, the linear correlation coefficient, R, values were
317 calculated for the same station cases mentioned under section 3.1. The obtained R values are given in **Table**
318 **3**. Based on these R values, ranks were assigned to indicate the strength of statistical homogeneity for the
319 tested cases. Lower the rank number, the higher the strength of statistical homogeneity. Among every two
320 stations tested, SS-20 versus SS-21, SS-26 versus SS-27 and SS-31 versus SS-32 produced the highest R
321 values (0.50-0.58) indicating highest chance of statistical homogeneity. SS-5 versus SS-6, SS-10 versus
322 SS-11, SS-14 versus SS-15, SS-15 versus SS-16, SS-18 versus SS-19, SS-19 versus SS-20, SS-25 versus
323 SS-26, SS-27 versus SS-28, SS-28 versus SS-29 and SS-30 versus SS-31 produced moderate R values
324 (0.32-0.48) indicating moderate chance of statistical homogeneity. SS-2 versus SS-3, SS-3 versus SS-4,
325 SS-4 versus SS-5, SS-6 versus SS-7, SS-9 versus SS-10, SS-11 versus SS12, SS-16 versus SS-17, SS-22
326 versus SS23, and SS-24 versus SS-25 produced low R values (0.20-0.25) indicating low chance of statistical
327 homogeneity. SS-1 versus SS-2, SS-7 versus SS-8, SS-8 versus SS-9, SS-12 versus SS-13, SS-13 versus

328 SS-14, SS-17 versus SS-18 and SS-21 versus SS-22 gave the lowest R values (less than 0.20) indicating
329 poor chance of statistical homogeneity. Note that the selection of demarcation boundaries for high, moderate,
330 low, and poor was done subjectively based on the orientation distributions observed on the 324-patch
331 network and the magnitude of R value. **Fig. 8** shows the pole plot comparisons obtained on the 324-patch
332 network for the cases of SS-7 versus SS-8, SS-2 versus SS-3, SS-5 versus SS-6 and SS-26 versus SS-27 on
333 R analysis as typical examples for poor, low, moderate, and high R values, respectively. Rankings seem to
334 agree very well with the visual orientation distributions seen on 324-patch networks.

335

336 **Fig. 8.** Pole plot comparisons on the 324-patch network for R analysis as typical examples for poor, low,
337 moderate, and high R values: (a) SS-7 versus SS-8 (poor $R = -0.04$), (b) SS-2 versus SS-3 (low $R = 0.24$),
338 (c) SS-5 versus SS-6 (moderate $R = 0.38$) and (d) SS-26 versus SS-27 (high $R = 0.58$).

339

340 Orientation data of locations that indicated existence of either high or moderate R values among
341 adjacent stations were then combined as follows to calculate R values for the combined stations: (a) SS-14
342 & 15 versus SS-16, (b) SS-18 versus SS-19 & 20, (c) SS-18 & 19 versus SS-20 & 21, (d) SS-26 & 27 versus
343 SS-28, (e) SS-26 & 27 versus SS-28 & 29 and (f) SS-30 versus SS-31 & 32. The obtained R values are
344 given in **Table 3**. Based on these R values, rankings were assigned based on the rankings 1-31 made earlier
345 using the first 32 cases tested. If two numbers are used for the ranking, it means that the ranking is in-
346 between the two numbers. SS-18 versus SS-19 and 20, and SS-18 and SS-19 versus SS-20 and SS-21
347 produced the same R value of 0.57, which is higher than that obtained for SS-18 versus SS-19 ($R = 0.37$),
348 SS-19 versus SS-20 ($R = 0.45$) and SS-20 versus SS-21 ($R = 0.50$). R values obtained for SS-26 versus SS-
349 27, SS-27 versus SS-28 and SS-28 versus SS-29 were 0.58, 0.48 and 0.35, respectively. R values obtained
350 for SS-26 and SS-27 versus SS-28, and SS-26 and SS-27 versus SS-28 and SS-29 were 0.50 and 0.55,
351 respectively. That means for these cases, the combined stations have produced R values which are in-
352 between the R values obtained for the uncombined cases. SS-14 and SS- 15 versus SS-16 produced a R
353 value of 0.30, which is lower than that obtained for the uncombined stations SS-14 versus SS-15 ($R = 0.41$)
354 and SS-15 versus SS-16 ($R = 0.33$). SS-30 versus SS-31 and SS-32 produced a R value of 0.38, which is
355 lower than that obtained for the uncombined stations SS-30 versus SS-31 ($R = 0.43$) and SS-31 versus SS-
356 32 ($R = 0.56$). The afore-mentioned examples show that different scenarios are possible by combining the
357 orientation data of adjacent stations. Out of all the combined stations compared, the stations SS-18 through
358 SS- 21 seem to indicate the highest level of statistical homogeneity. Stations SS-26 through SS-29 indicate
359 high level of statistical homogeneity. Stations SS-14 through SS-16 and SS-30 through SS-32 show
360 moderate level of statistical homogeneity.

361

362 4.3. Use of contingency table analysis to assign quantitative ranks of statistical homogeneity

363 Using the procedure explained in section 3.3, the average p values were calculated for the same station
364 cases mentioned under section 3.1. The obtained average p values are given in **Table 3**. Based on these
365 average p values, ranks were assigned to indicate the strength of statistical homogeneity. Lower the rank
366 number, the higher the strength of statistical homogeneity. Among every two sites tested, SS-26 versus SS-
367 27, SS-20 versus SS-21, SS-10 versus SS-11 and SS-19 versus SS-20 produced the highest average p values
368 (0.37-0.66) indicating high chance of statistical homogeneity. SS-27 versus SS-28, SS-31 versus SS-32,
369 SS-5 versus SS-6, SS-30 versus SS-31, SS-16 versus SS-17, SS-14 versus SS 15, SS-15 versus SS-16 and
370 SS-18 versus SS-19 produced moderate average p values (0.16-0.33) indicating moderate chance of
371 statistical homogeneity. SS-22 versus SS-23, SS-4 versus SS-5, SS-28-versus SS-29, SS-11 versus SS-12
372 and SS-6 versus SS-7 produced low average p values (0.05-0.16) indicating low chance of statistical
373 homogeneity. The rest of every two stations tested gave the lowest average p values (0-0.05) indicating
374 poor chance of statistical homogeneity. Note that the demarcation boundary between poor and low
375 statistical homogeneity was selected based on the minimum significance level statisticians use for
376 acceptance of the null hypothesis of statistical homogeneity. Demarcation boundaries between low and
377 moderate, and moderate and high were selected subjectively based on the obtained average p values and
378 the visual observation of orientation distributions on the 68-patch network.

379
380 Orientation data of locations that indicated existence of either high or moderate average p values among
381 adjacent stations were then combined as follows to calculate average p values for the combined stations:
382 (a) SS-14 & 15 versus SS-16, (b) SS-18 versus SS-19 & SS-20, (c) SS-18 & SS-19 versus SS-20 & SS-21,
383 (d) SS-26 & SS-27 versus SS-28, (e) SS-26 & 27 versus SS-28 & SS-29 and (f) SS-30 versus SS-31 & SS-
384 32. The obtained p values are given in **Table 3**. Based on these average p values, rankings were assigned
385 based on the rankings 1-31 made earlier using the first 32 cases tested. If two numbers are used for the
386 ranking, it means that the ranking is in-between the two numbers. SS-18 versus SS-19 and SS-20 produced
387 a p value of 0.45, which is higher than that obtained for SS-18 versus SS-19 ($p = 0.16$) and SS-19 versus
388 SS-20 ($p = 0.38$). For this case, the statistical homogeneity has increased by combining the data of adjacent
389 stations. On the other hand, for all other combined station cases the obtained average p values were
390 significantly less than that obtained for the corresponding uncombined two station cases (see **Table 3**). This
391 indicates that the statistical homogeneity has decreased for these cases by combining the data from adjacent
392 stations. Based on the contingency table analysis results, out of all the combined stations compared, the
393 stations SS-18 through SS- 20 seem to indicate high level of statistical homogeneity. Stations SS-14 through
394 SS-16 indicate low level of statistical homogeneity. SS-26 through SS-29 and SS-30 through SS-32 show
395 poor level of statistical homogeneity. Note that the low and poor level statistical homogeneity results

396 obtained for the combined stations through contingency table analysis results are very different to the results
397 obtained through R calculations. This is currently a puzzle. Note that the number of data increases when
398 we combine the stations. The afore-mentioned discrepancy between the R calculation results and
399 contingency analysis results might have resulted from increase of discrimination power of the contingency
400 table analysis procedure with the increase of the number of data of the tested cases.

401

402 **5. Conclusions**

403 As shown in this paper, no solid objective method is available at present in the engineering geology or
404 rock mechanics literature to firmly demarcate between the statistical homogeneity and non-homogeneity of
405 rock masses. Each of the available methods can only be used to rank rock masses with respect to the strength
406 of statistical homogeneity. Three different methods, (a) polar equal area pole plot comparison, (b) linear
407 correlation coefficient analysis and (c) contingency table analysis, were applied in this paper to evaluate
408 the statistical homogeneity of rock masses that exist along a highway corridor in Vietnam. The quantitative
409 ranks obtained from the linear correlation analysis and the contingency table analysis have some similarities
410 as well as some differences (see **Table 3**). Analysis performed for every two sites have provided
411 quantitative rankings within 5 for 21 out of 31 tested cases, which is a significant similarity. For 2 out of
412 31 cases, the ranking difference is more than 12, which is a major difference. For 8 out of 32 cases, the
413 ranking difference is between 6 and 8, which is a significant difference. For the calculations conducted on
414 the combined stations through the two quantitative methods, only one case out of six has given a similar
415 result; the rest of the five cases show drastic differences. Comparison of the results from the two quantitative
416 methods clearly show that SS-26 versus SS-27 and SS-20 versus SS-21 indicate high chance of rock mass
417 statistical homogeneity; SS-10 versus SS-11, SS-19 versus SS-20, SS-27 versus SS-28 and SS-31 versus
418 SS-32 indicate medium to high statistical homogeneity. Comparison of the results from the two quantitative
419 methods for SS-5 versus SS-6, SS-30 versus SS-31, SS-14 versus SS-15, SS-15 versus SS-16, SS-18 versus
420 SS-19, SS-28 versus SS-29 and SS-16 versus SS-17 indicate medium level of rock mass statistical
421 homogeneity. The compared results from the two quantitative methods for SS-22 versus SS-23, SS-4 versus
422 SS-5, SS-11 versus SS-12 and SS-6 versus SS-7 show low level of rock mass statistical homogeneity.
423 Comparison of the results from the two quantitative methods for the rest of the cases show poor level of
424 statistical homogeneity. According to the quantitative ranking results, SS-26 versus SS-27 can be declared
425 as the stations that show the highest chance of rock mass statistical homogeneity; SS-7 versus SS-8 can be
426 declared as the stations that show the poorest chance of rock mass statistical homogeneity. For majority of
427 the cases, the conclusions obtained through comparison of the rankings obtained from the two quantitative
428 methods seem to agree reasonably well with the subjective qualitative rankings obtained through visual
429 comparison of polar equal area pole plots of the orientation data. The study clearly illustrates that

430 application of different methodologies can produce different ranking orders with respect to the strength of
431 statistical homogeneity of rock masses. This finding clearly indicates that decision making on statistical
432 homogeneity of rock masses is a difficult task using results coming from just one method. This case study
433 indicates that comparison of results coming from several methods allow one to make better judgement on
434 rock mass statistical homogeneity.

435

436 **References**

437 Goodman, R.E., 1976. Methods of geological Engineering in discontinuous rocks. West Publishing Co, St
438 Paul, Minn.

439 Hoek, E., Bray, J.W., 2004. Rock slope Engineering. Taylor & Francis Group, London, and New York,
440 431p.

441 Kasatskin, S.A., Golozubov, V.V., Phung, V.P., Le, D.A., 2014. Evidence of Cenozoic Strike-Slip
442 Dislocations of the Red River fault system in Paleozoic Carbonate Strata of Cat Ba Island (northern
443 Vietnam) Russian Journal of Pacific Geology. 8(3), 163-176.

444 Kulatilake, P.H.S.W., Chen, J., Teng, J., Shufang, X., Pan, G., 1996. Discontinuity geometry
445 characterization for the rock mass around a tunnel close to the permanent ship-lock area of the Three Gorges
446 Dam site in China. Int. J. Rock Mech. and Min. Sci. 33, 255-277.

447 Kulatilake, P.H.S.W., Fiedler, R., Panda, B.B., 1997. Box fractal dimension as a measure of statistical
448 homogeneity of jointed rock masses. Engineering Geology. 48(3-4), 217-230.

449 Kulatilake, P.H.S.W., Wathugala, D.N., Poulton, M., Stephansson, O., 1990. Analysis of structural
450 homogeneity of rock masses. Engineering Geology. 29(3), 195-211.

451 Lancaster, H.O., 1969. The chi-squared distribution. Wiley, New York, 175p.

452 Li, Y., Wang, Q., Chen, J., Song, S., Ruan, Y., Zhang, Q., 2015. A multivariate technique for evaluating
453 the statistical homogeneity of jointed rock masses. Rock Mechanics and Rock Engineering. 48(5), 1821-
454 1831.

455 Mahtab, M.A., Yegulalp, T.M., 1984 A similarity test in grouping data in rock mechanics. Proc 25th US
456 Symposium in Rock Mechanics. New York, 495-502.

457 Martin, M.W., Tannant, D.D., 2004. A technique for identifying structural domain boundaries at the EKATI
458 Diamond Mine. Engineering Geology. 74, 247-264.

459 Miller, S.M., 1983 A statistical method to evaluate homogeneity of structural populations. Math Geol. 15,
460 317-328.

461 Nguyen, T.Y., 1991. Main features of modern geodynamic in the North Vietnam. Geology-Resource,
462 National Centre for Natural Science and Technology Institute of Geology, 7-10. In Vietnamese.

463 Nguyen, K.Q., Dinh, T.T., Tran, V.T., Dao, D.T., Le, V.C., Nguyen, D.D., Nguyen, T.V., Nguyen, V.H.,
464 Pham, V.H., Phan, C.T., Tong, D.T., Tran, T.T., Trinh, D., Vu, K., 2000. Geological and Mineral Resources

- 465 Map of Viet Nam on 1:200.000: Backan (F-48-XVI). Department of Geology and Minerals of Vietnam,
466 Hanoi.
- 467 Phi, N.Q., Gi, H.S., Thanh, P.T., Phuong, N., 2012. Structural domain identification by fracture orientation
468 and fracture density in rock mass. *International Journal of Geoinformatics*. 8, 35-40.
- 469 Phung, V.P., Nguyen, T.Y., Vu, V.C., 1996. Geodynamic situation in neotectonic and recent period on
470 territory of Vietnam. *Geology-Resource*, National Centre for Natural Science and Technology institute of
471 geology. I, 101-111. In Vietnamese.
- 472 Rocscience, Inc., 2019. *Dips v7.0 graphical and statistical analysis of orientation data*. Toronto, Ontario,
473 Canada
- 474 Song, S., Sun, F., Zhang, W., Chen, J., Xu P., Niu, C., Cao, C., Zhan, J., 2018. Identification of structural
475 domains by considering multiple discontinuity characteristics: a case study of the Songta Dam. *Bulletin of*
476 *Engineering Geology and the Environment*. 77(4), 1589-1598.
- 477 Thanh, T.P., Gi, H.S., Phi, N.Q., 2015. Delineate structural boundary from fracture correlation coefficients.
478 *Vietrock 2015 an ISRM specialized conference*, Hanoi, 243-251.
- 479 Thanh, P.T., 2019. Tectonic Activity Phases of Cenozoic Period in Xuat Hoa Area, Bac Kan Province,
480 Northeast Region, Vietnam. *Indonesian Journal on Geoscience*. 6(3), 311-325.

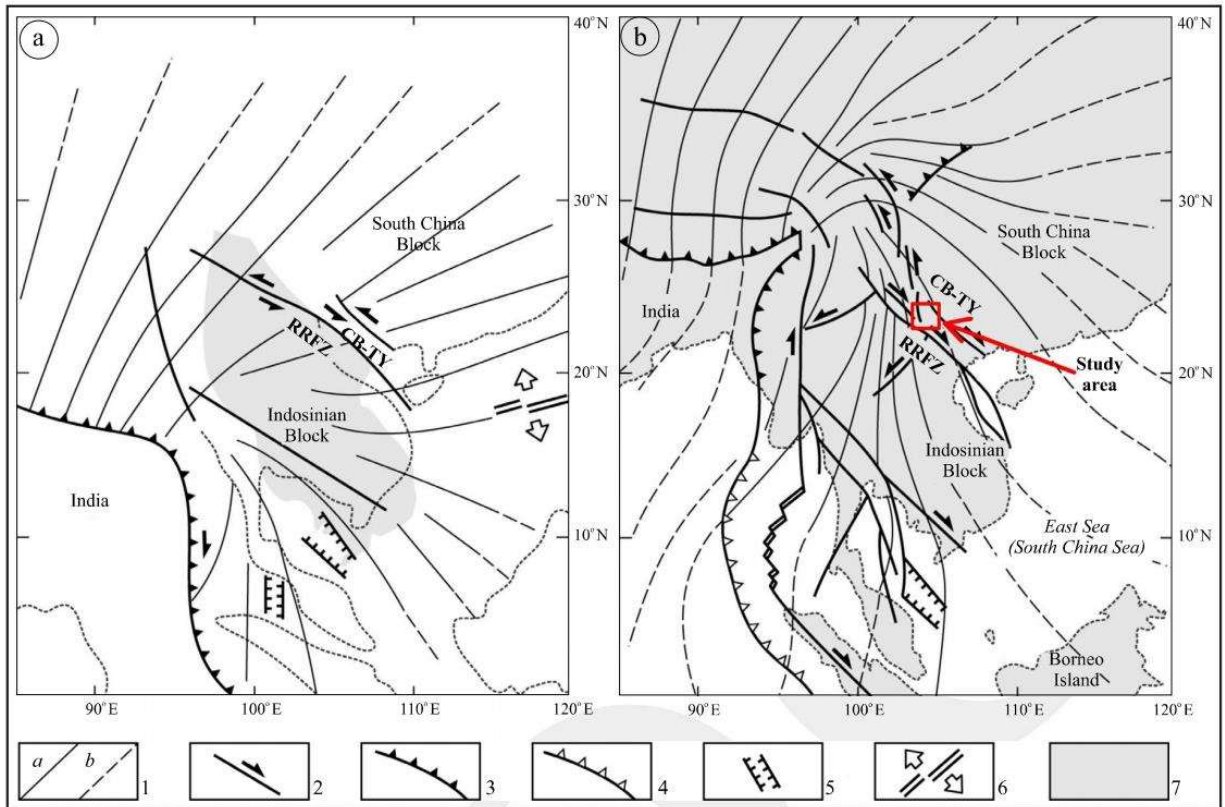


Fig. 1. Trajectories of the maximum compressive stress within the Indochina Peninsula during (a) the Oligocene and (b) at the present time (Kasatkin et al., 2014). The legend on the map in Figure 1 is described as follows: 1. trajectories of the maximum compressive stress: (a) which are directly related to the Indo-Eurasian plate collision and (b) its far-field effects; (2) faults and directions of displacement (arrows); (3) zone of continental collision; (4) subduction zone; (5) extension structures; (6) spreading zones; (7) current position of the land; Red River Fault System (RRFS); Cao Bang - Tien Yen fault (CB - TY).

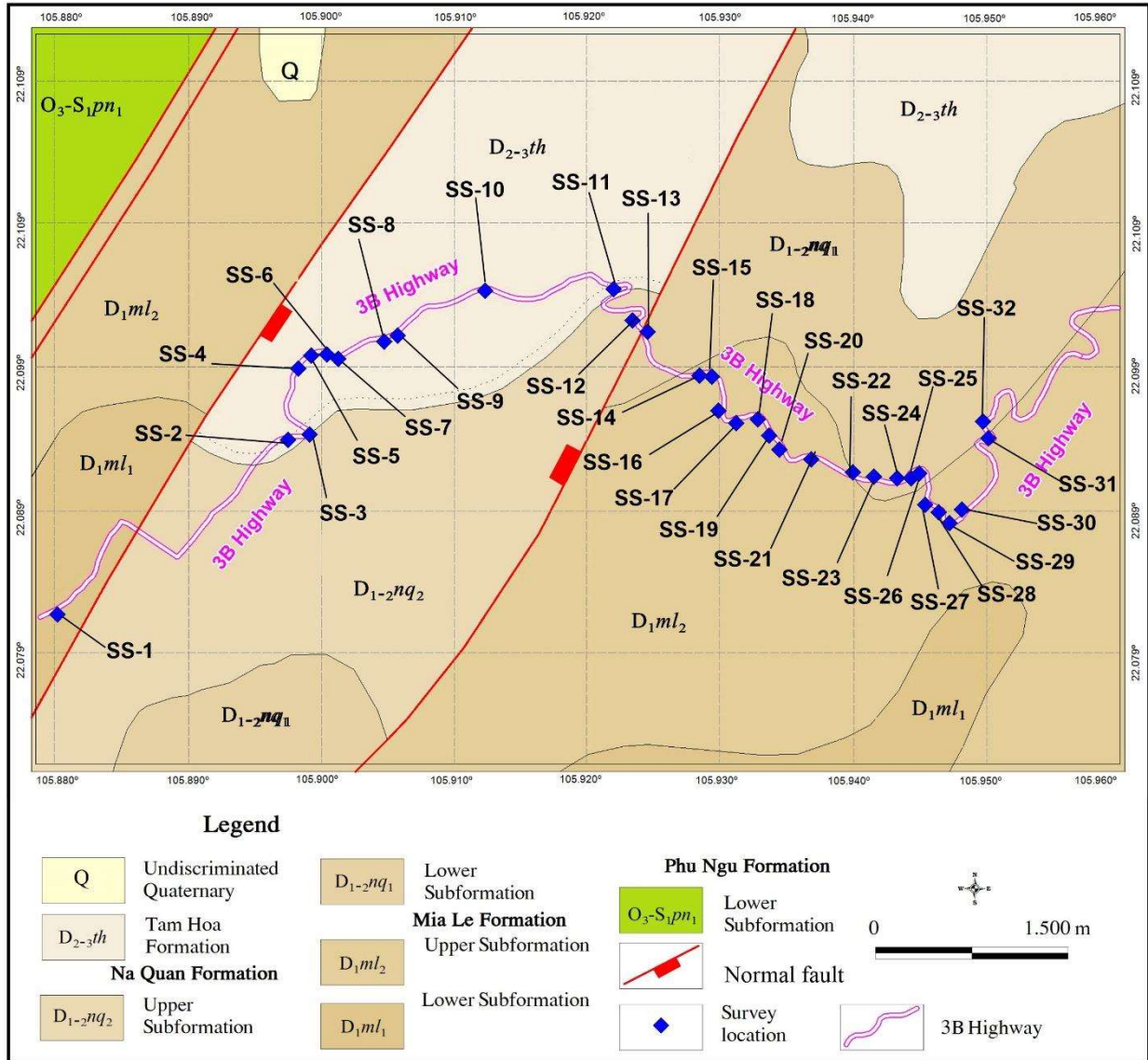
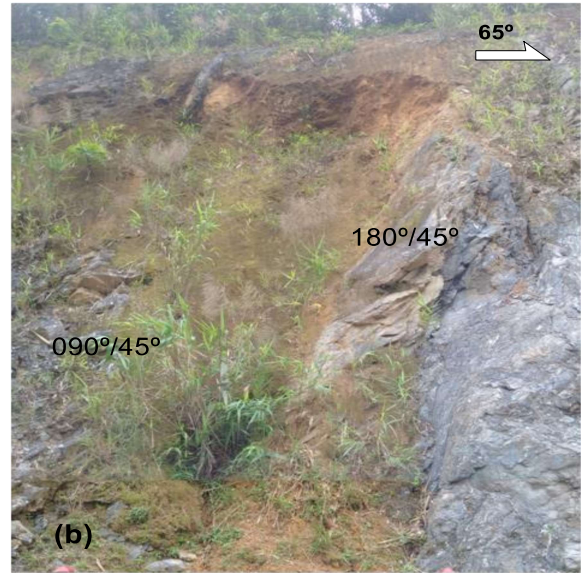


Fig. 2. Geological map of the study area showing the survey stations (SS) used to obtain orientation data.



(a) Tam Hoa (D_{2-3th}) formation at SS-11



(b) Na Quan Formation (D_{1-2nq2}) at site SS-12



(c) Mia Le Formation (D_{1ml2}) at site SS-19



(d) Mia Le Formation (D_{1ml2}) at site SS-27

Fig. 3. Typical rock masses that exist at a few survey stations.

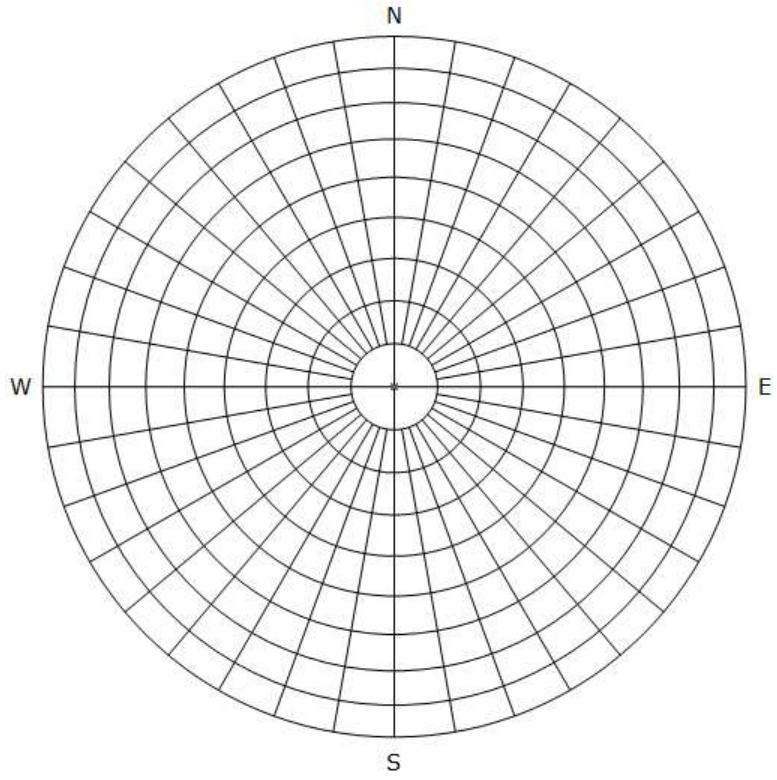


Fig. 4. The constructed 324-patch polar equal area net.

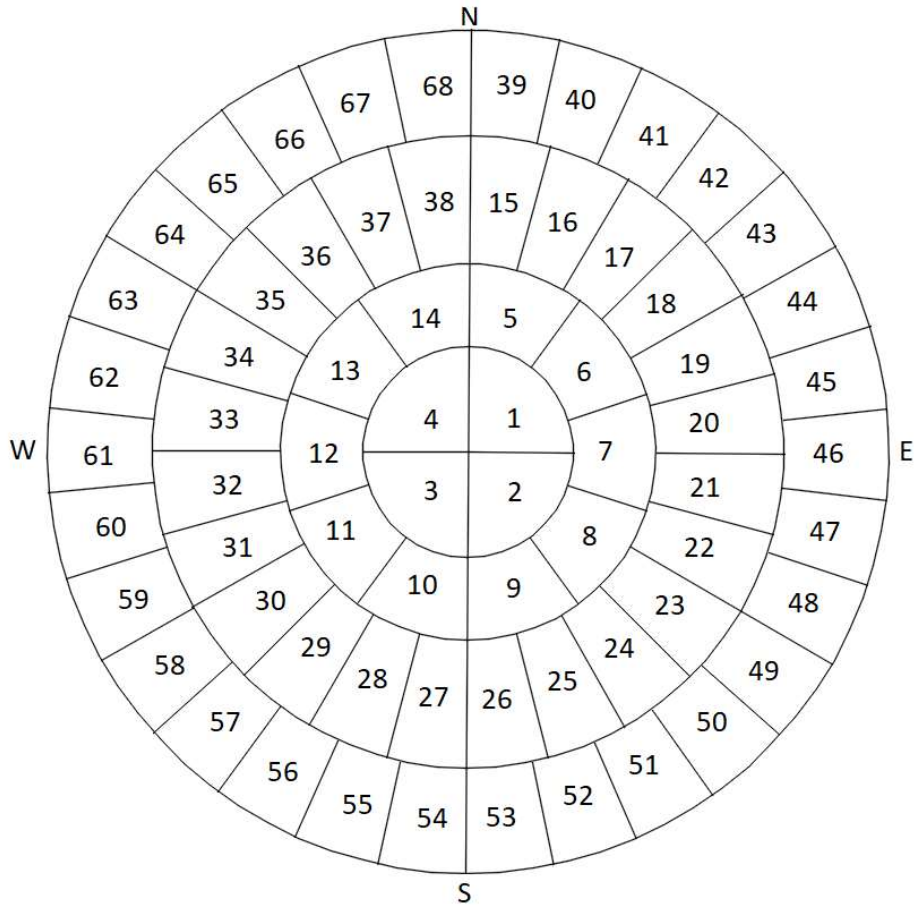
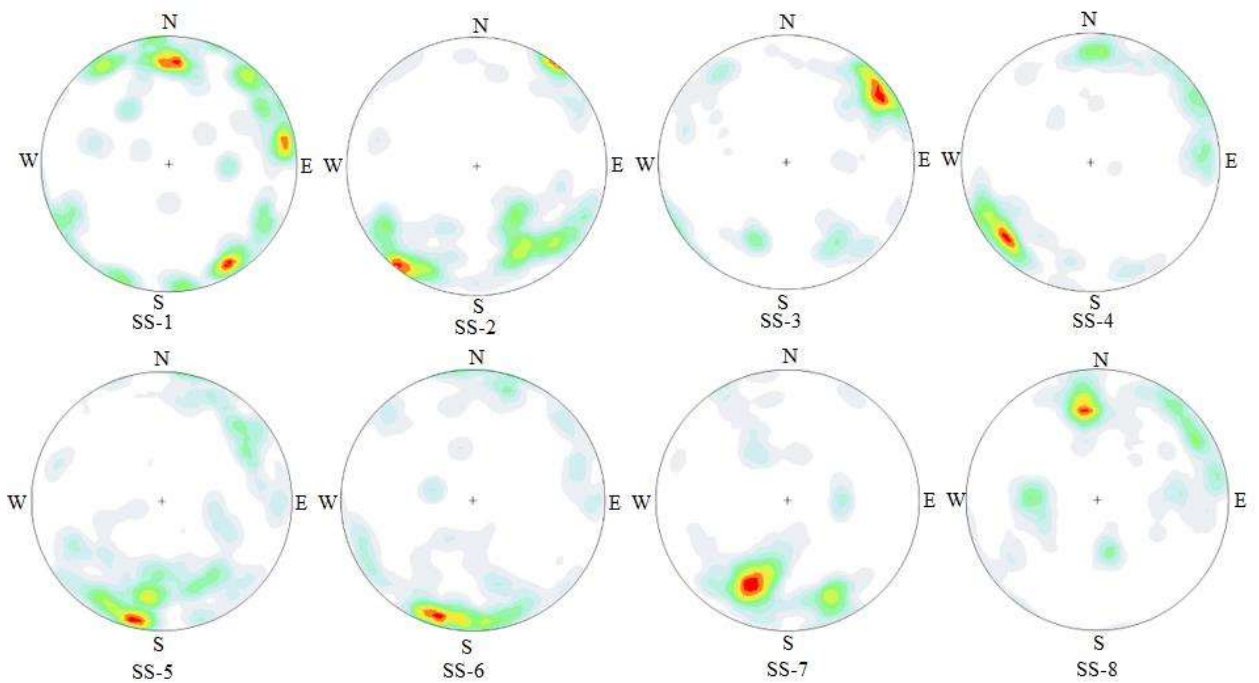
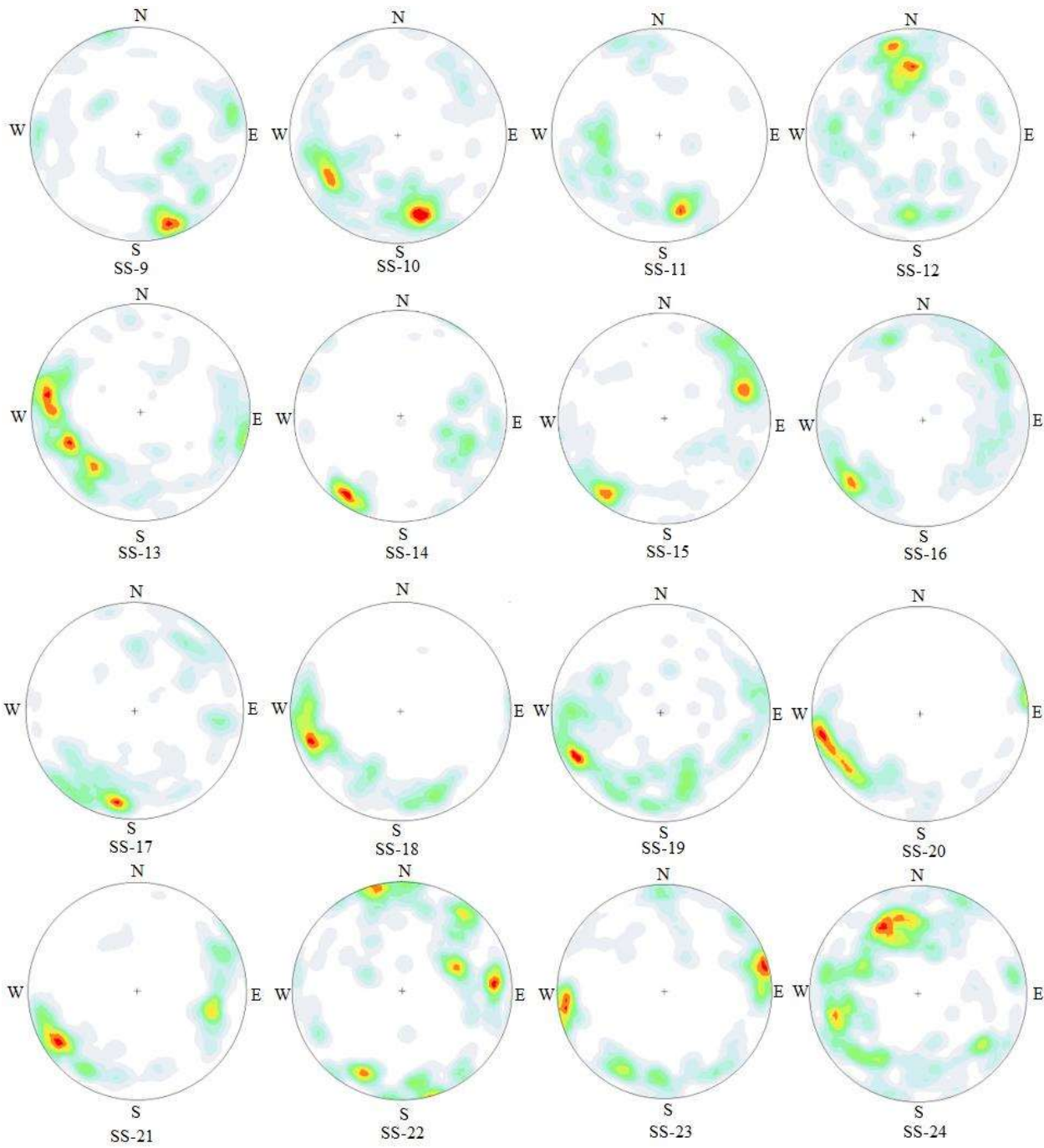


Fig. 5. The designed 68-cell patch network used for the contingency table analysis.





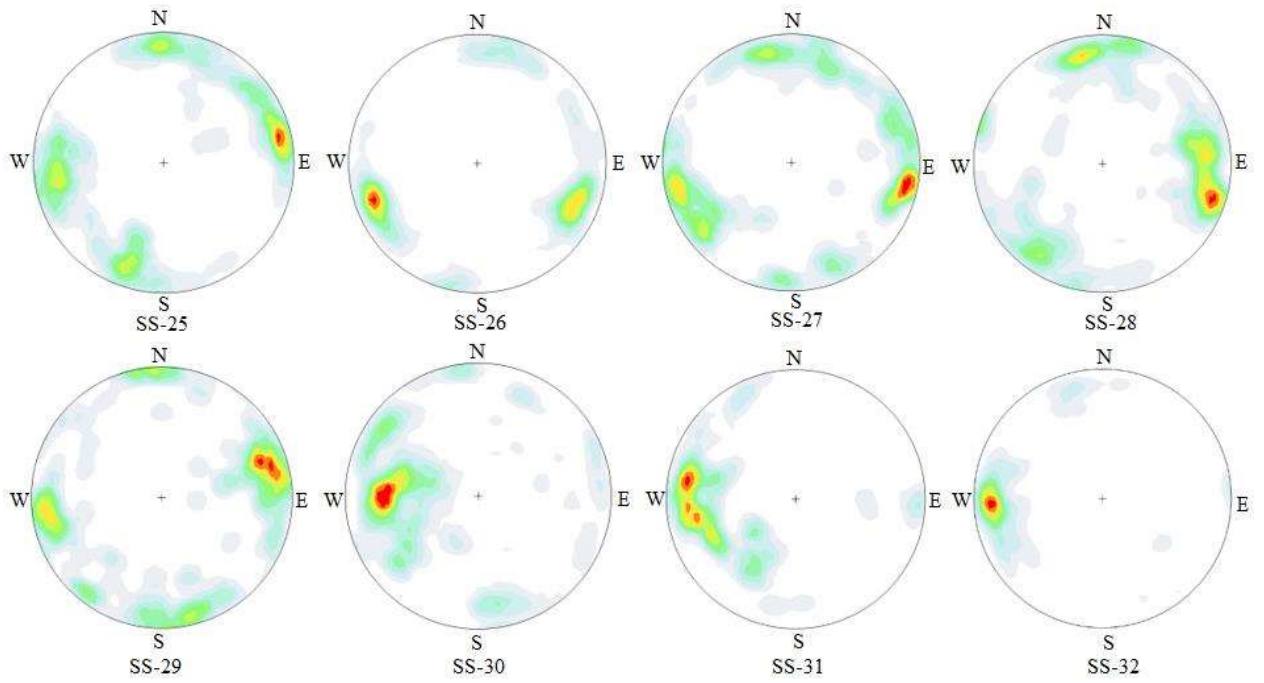


Fig. 6. Lower hemispherical polar equal area pole plots of the 32 stations located along the highway corridor.

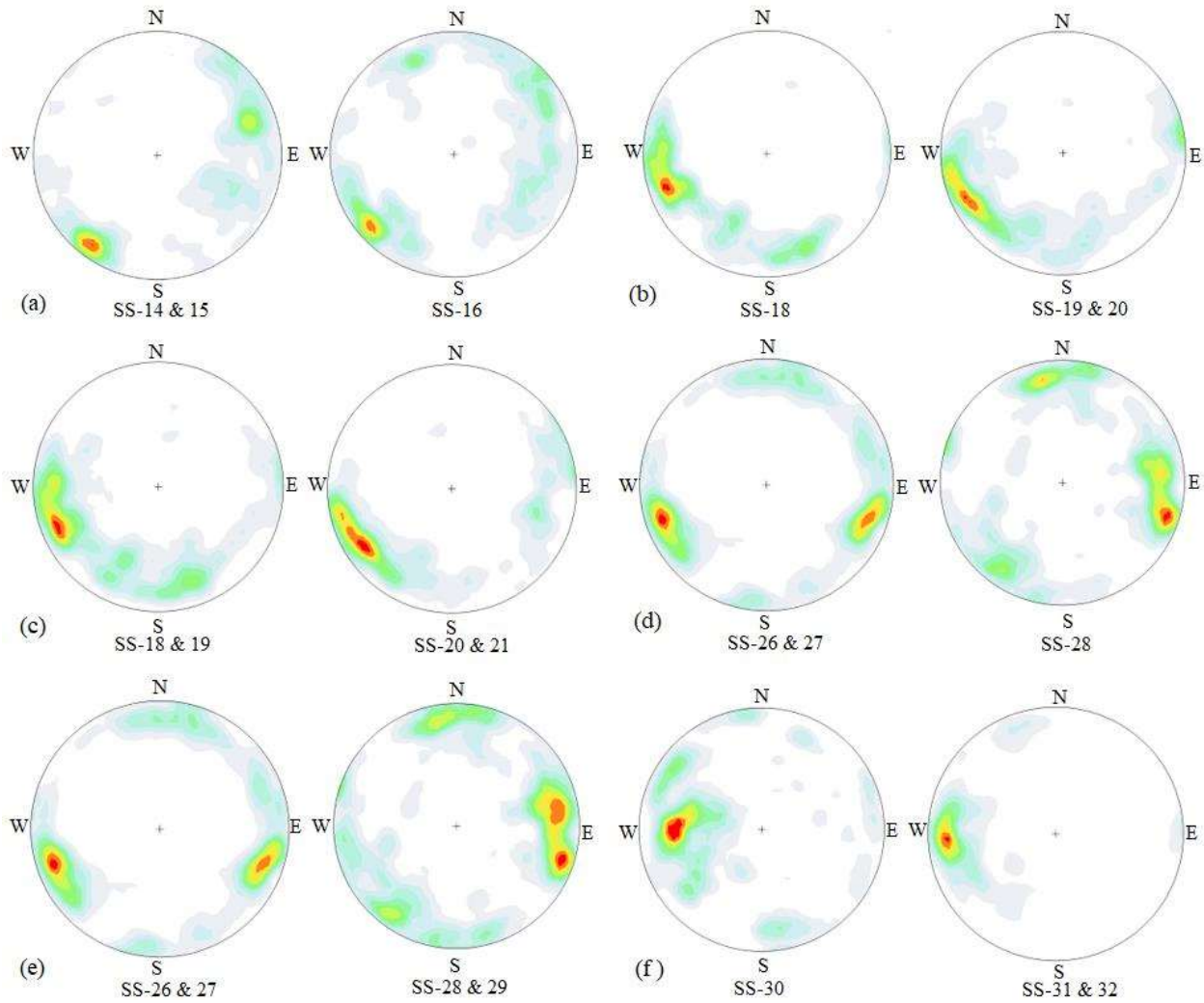


Fig. 7. Lower hemispherical polar equal area pole plot comparisons made for the combined stations: (a) SS-14 & 15 versus SS-16 (L level), (b) SS-18 versus SS-19 & 20 (H level), (c) SS-18 & 19 versus SS-20 & 21 (M level), (d) SS-26 & 27 versus SS-28 (L level), (e) SS-26 & 27 versus SS-28 & 29 (L level) and (f) SS-30 versus SS-31 & 32 (L level).

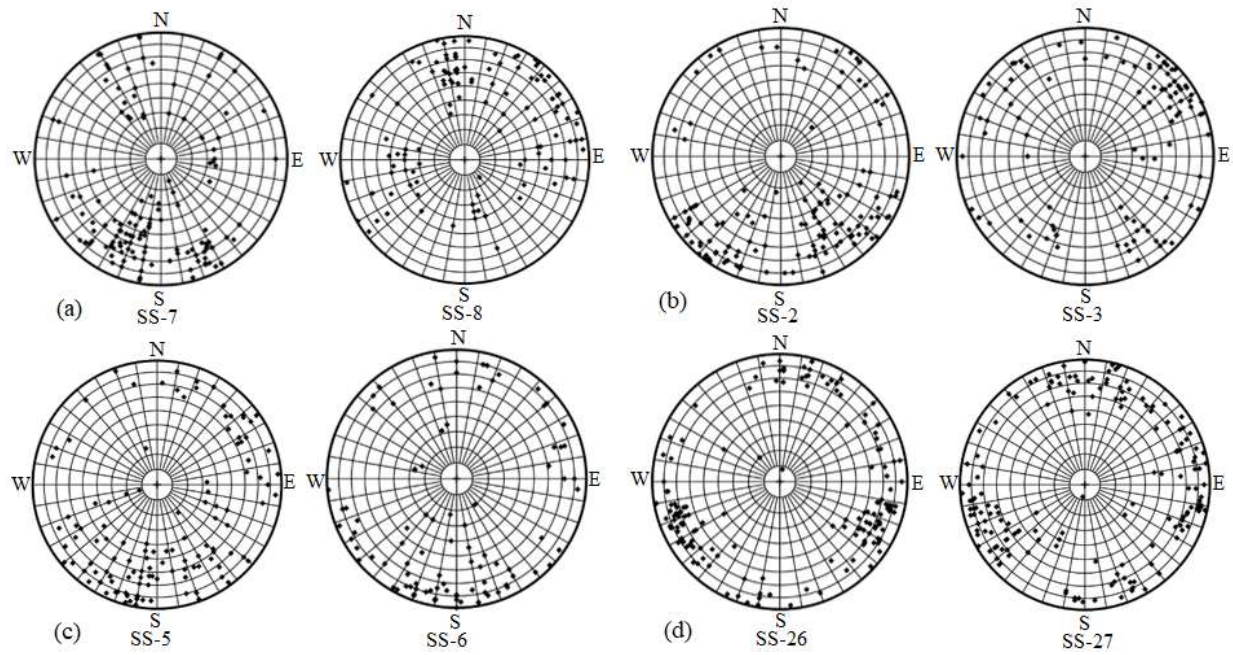


Fig. 8. Pole plot comparisons on the 324-patch network for R analysis as typical examples for poor, low, moderate, and high R values: (a) SS-7 versus SS-8 (poor $R = -0.04$), (b) SS-2 versus SS-3 (low $R = 0.24$), (c) SS-5 versus SS-6 (moderate $R = 0.38$) and (d) SS-26 versus SS-27 (high $R = 0.58$).

Table 1 Survey station locations, geologic age and the number of discontinuity orientation data collected for the various stations.

Survey station number	Longitude	Latitude	Geological age	Number of discontinuities
SS-1	105°52'56.8"	22°04'46.5"	D₁ml₁	73
SS-2	105°53'52.2"	22°5'37.6";	D₂₋₃th	127
SS-3	105°53'53.7"	22°05'09.2"	D₂₋₃th	103
SS-4	105°53'56.2"	22°05'55.8"	D₂₋₃th	116
SS-5	105°53'59.8"	22°05'59.1"	D₂₋₃th	122
SS-6	105°54'03.9"	22°05'59.4"	D₂₋₃th	96
SS-7	105°54'07.0"	22°05'58.2"	D₂₋₃th	137
SS-8	105°54'19.5"	22°06'02.5"	D₂₋₃th	119
SS-9	105°54'23.1"	22°06'03.9"	D₂₋₃th	105
SS-10	105°54'46.9"	22°06'15.1"	D₂₋₃th	188
SS-11	105°52'21.8"	22°06'15.6"	D₂₋₃th	136
SS-12	105°55'28.2"	22°06'09.3"	D₁₋₂nq₂	113
SS-13	105°55'30.8"	22°06'04.6"	D₁₋₂nq₂	135
SS-14	105°55'44.7"	22°05'53.4"	D₁ml₂	71
SS-15	105°55'48.1"	22°05'53.1"	D₁ml₂	167
SS-16	105°55'49.8"	22°05'44.6"	D₁ml₂	165
SS-17	105°55'54.6"	22°05'41.4"	D₁ml₂	141
SS-18	105°56'00.5"	22°05'42.3"	D₁ml₂	104
SS-19	105°56'03.5"	22°05'38.2"	D₁ml₂	120
SS-20	105°56'06.2"	22°05'34.7"	D₁ml₂	103
SS-21	105°56'14.8"	22°05'32.1"	D₁ml₂	119
SS-22	105°56'26.0"	22°05'28.8"	D₁₋₂nq₁	99
SS-23	105°56'31.8"	22°05'27.7"	D₁₋₂nq₁	128
SS-24	105°56'38.1"	22°05'27.2"	D₁₋₂nq₁	152
SS-25	105°56'41.8"	22°05'27.3"	D₁₋₂nq₁	155
SS-26	105°56'44.1"	22°05'24.7"	D₁₋₂nq₁	158
SS-27	105°56'45.6"	22°05'20.6"	D₁ml₂	172
SS-28	105°56'49.2"	22°05'18.6"	D₁ml₂	215
SS-29	105°56'52.1"	22°05'15.9"	D₁ml₂	102
SS-30	105°56'55.5"	22°05'19.3"	D₁ml₂	205
SS-31	105°57'02.8"	22°05'37.2"	D₁ml₂	103
SS-32	105°57'01.3"	22°05'41.4"	D₁₋₂nq₁	140

Table 2 Contingency table for analyzing the polar equal area plots of different stations (after Miller, 1983).

Columns				
Rows	Patch 1	Patch 2	Patch 3	Row totals
Equal area plot 1	f_{11}	f_{12}	f_{1c}	R_1
Equal area plot 2	f_{21}	f_{22}	f_{2c}	R_2
·	-	-	-	-
·	-	-	-	-
·	-	-	-	-
Equal area plot r	f_{r1}	f_{r2}	f_{rc}	R_r
Column totals	C_1	C_2	C_c	N

f_{ij} = number of discontinuity poles observed in cell ij ; N = total number of discontinuity poles from all equal area plots.

Table 3 Summary results for strength of statistical homogeneity based on different methods.

Tested combinations	Avg. max. sig. value based on normal dist.	Avg. max. sig. value based on chi-square dist.	Avg. max. sig. value based on chi-square and normal dists.	Rank based on contingency table analysis results	Rank based on R value	R value	Qualitative rank based on equal area and pole plot comparisons
SS1-SS2	0.0216	0.0351	0.0283	18	23	0.18	P
SS2-SS3	0.003	0.0112	0.0071	23	16	0.24	L-P
SS3-SS4	0.0007	0.0054	0.0062	24	21	0.21	L-P
SS4-SS5	0.1711	0.1222	0.1467	14	20	0.22	L
SS5-SS6	0.3214	0.184	0.2527	7	9	0.38	M
SS6-SS7	0.1199	0.0921	0.106	17	22	0.20	L
SS7-SS8	0	0	0	31	31	-0.04	P
SS8-SS9	0	0.0002	0.0001	30	29	0.08	P
SS9-SS10	0.0078	0.0156	0.0117	20	14	0.25	L-P
SS10-SS11	0.5002	0.2708	0.3855	3	5	0.47	M
SS11-SS12	0.1508	0.114	0.1324	16	16	0.24	L
SS12-SS13	0.005	0.0131	0.0091	22	30	0.04	P
SS13-SS 14	0	0.0008	0.0004	27	28	0.11	P
SS14-SS15	0.2076	0.1365	0.1721	10	8	0.41	M
SS15-SS16	0.1965	0.14	0.1683	11	12	0.33	M
SS16-SS17	0.2172	0.1383	0.1778	9	16	0.24	L
SS17-SS18	0	0.0003	0.0002	29	27	0.15	P
SS18-SS19	0.1909	0.1353	0.1631	12	10	0.37	M
SS19-SS20	0.4857	0.2701	0.3779	4	6	0.45	M
SS20-SS21	0.5792	0.5048	0.542	2	3	0.50	H
SS21-SS22	0.0046	0.0168	0.0107	21	23	0.18	P
SS22-SS23	0.1959	0.1214	0.1586	13	19	0.23	L
SS23-SS24	0.0182	0.0088	0.0135	19	23	0.18	P
SS24-SS25	0	0.0005	0.0003	28	14	0.25	P
SS25-SS26	0	0.001	0.0005	26	13	0.32	L-P
SS26-SS27	0.8679	0.4426	0.6553	1	1	0.58	H
SS27-SS28	0.4129	0.2343	0.3236	5	4	0.48	H
SS28-SS29	0.1607	0.1117	0.1362	15	11	0.35	M
SS29-SS30	0.0002	0.002	0.0011	25	26	0.16	P
SS30-SS31	0.2147	0.1428	0.1788	8	7	0.43	M
SS31-SS32	0.3998	0.2297	0.3148	6	2	0.56	H
SS14&15-SS16	0.0526	0.0601	0.0564	17-18	13-14	0.30	L
SS18-SS19&20	0.5658	0.3348	0.4503	2-3	1-2	0.57	H
SS18&19-SS20&21	0.0517	0.0527	0.0522	17-18	1-2	0.57	M
SS26&27-SS28	0.0036	0.0095	0.0066	23-24	3	0.50	L
SS26&27-SS28&29	0.001	0.0051	0.0031	24-25	2-3	0.55	L
SS30-SS31&32	0.0006	0.0096	0.0051	24-25	9	0.38	L

Conflict of interest/competing interest

The authors declare that they have no conflict of interest.